

## Solutions to HW #6

### Problem 1

$$(a) \quad R_{NO} - G_o = H_o + \lambda E_o = 500 \frac{W}{m^2}$$

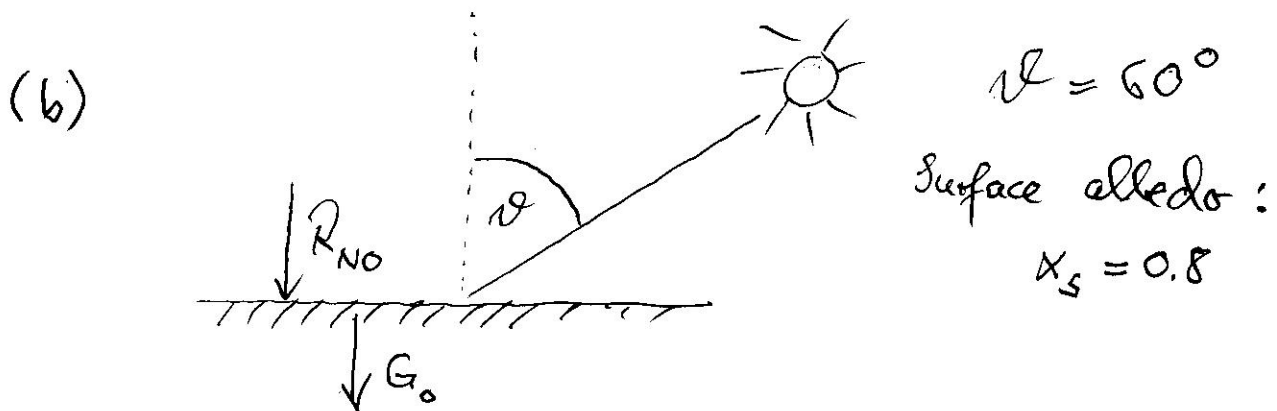
$$\text{Bowen ratio: } \beta = \frac{H_o}{\lambda E_o}$$

$$\beta = 3 \Rightarrow H_o = 3 \lambda E_o \Rightarrow H_o \gg \lambda E_o \Rightarrow \text{dry conditions}$$

$$\Rightarrow 500 \frac{W}{m^2} = 3 \lambda E_o + \lambda E_o = 4 \lambda E_o$$

$$\Rightarrow \lambda E_o = \frac{1}{4} \cdot 500 \frac{W}{m^2} = \underline{\underline{125 \frac{W}{m^2}}}$$

$$\Rightarrow H_o = 3 \lambda E_o = \underline{\underline{375 \frac{W}{m^2}}}$$



Net surface radiative flux:

$$R_{NO} = R_{SO} - \alpha_s R_{SO} + \epsilon_s R_{LO}^d - R_{LO}^u$$

Problem 1

 (cont'd)

(b) (cont'd)

Idealized conditions:

$$R_{SO} - \alpha_s R_{SO} = (1 - \alpha_s) S_0 \cdot \cos \theta = 0.2 \cdot 1366 \frac{W}{m^2} \cdot \underbrace{\cos 60^\circ}_{= \frac{1}{2}}$$

$$= 136.6 \frac{W}{m^2}$$

$$R_{LO}^d = 0, \quad R_{LO}^u = \epsilon_s \sigma T_s^4$$

$$= 1 \cdot 5.67 \cdot 10^{-8} \frac{W}{m^2 K^4} \cdot (273 K)^4$$

$$= 314.9 \frac{W}{m^2}$$

~ max. possible  
T for fresh snow

$G_0 = 0$  (let's assume fresh snow is an ideal isolator)

$$\Rightarrow R_{NO} - G_0 = 136.6 \frac{W}{m^2} - \sigma T_s^4$$

Obviously, the available energy  $R_{NO} - G_0$  decreases with increasing  $T_s$  and is negative unless it is very cold. The sign changes at  $T_s = \left( \frac{136.6 W/m^2}{\sigma} \right)^{\frac{1}{4}} = 221.5 K = \underline{\underline{-52^\circ C}}$

That is,  $R_{NO} - G_0$  reaches a maximum when  $T_s$  reaches its minimum.

### Problem 1 (cont'd)

(c) Same as (b), except  $R_{NO} - G_0$  is increased

$$\text{by } \epsilon_s R_{LO}^d = 1.567 \cdot 10^{-8} \frac{W}{m^2 K^4} \cdot [(273 - 50) K]^4$$

$$= 140 \frac{W}{m^2}$$

$$\Rightarrow R_{NO} - G_0 = \underbrace{137 \frac{W}{m^2} + 140 \frac{W}{m^2}}_{277 \frac{W}{m^2}} - \epsilon T_s^4$$

In this case, the equilibrium surface temperature would be  $T_s = -9^\circ C$ .

That is, if there is (fresh) snow on the ground, the available energy depends very sensitively on  $R_{LO}^d$  from clouds, even very high (and cold) clouds.

### Problem 2

Typo 1: Eq. (7.4) should read:  $\frac{D\underline{V}}{Dt} \downarrow = 2\underline{V} \otimes \underline{\Omega} - \frac{1}{S} \dots$ ,  
not  $\frac{D\underline{V}}{Dt} \downarrow + 2\underline{V} \otimes \underline{\Omega} \dots$

Typo 2: In Eq. (7.6), the first term in the parentheses in front of  $\underline{j}$  should be  $\frac{D\underline{r}}{Dt}$ , not  $\frac{D\underline{u}}{Dt}$ .

### Problem 3

Critical  $Ri$  is equal to  $\frac{1}{4} \Rightarrow \frac{1}{4} = \frac{N^2}{S_c^2}$

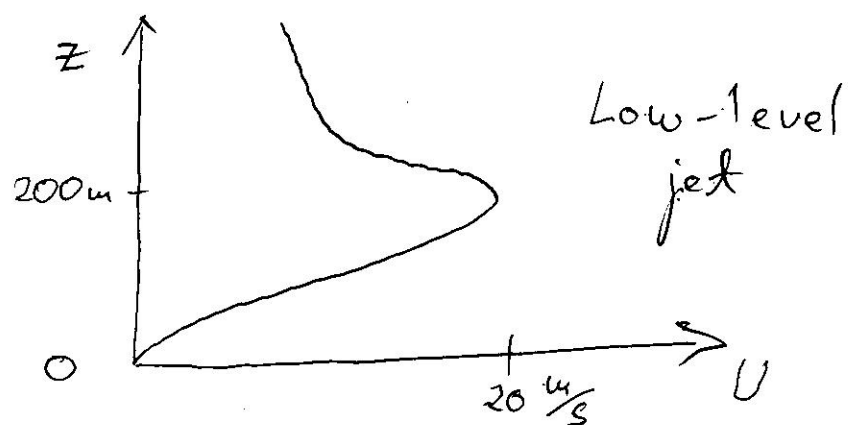
$\Rightarrow |S_c| = 2N$  is the critical shear.

$$\Rightarrow |S_c| = 2 \sqrt{\frac{g}{\theta} \frac{\partial \theta}{\partial z}} = 2 \sqrt{\frac{g}{T} \left( \frac{\partial T}{\partial z} + \frac{g}{\rho} \right)}$$

$$= 2 \sqrt{\frac{10 \frac{m}{s^2}}{300 K} \left( 10^{-2} \frac{K}{m} + 10^{-2} \frac{K}{m} \right)}$$

$$= 0.052 \frac{1}{s} = 0.052 \frac{m/s}{m} = \underline{\underline{5.2 \frac{m/s}{100 m}}}$$

low-level jets with  $U = 20 \frac{m}{s}$  at  $z = 200 m$  are quite common. Because  $U = 0$  at the ground (no-slip condition), shear rates of  $\frac{20 \frac{m}{s}}{200 m} = 0.1 \frac{1}{s}$  and more are to be expected. Such low-level jets are often turbulent, which is consistent with our result above.



Extra-credit problem

(a) Isothermal atmosphere:  $T(z) = T_0 = \text{const.}$

$$\text{Gas eq.: } \rho = \frac{p}{RT} \Rightarrow \rho = \frac{p}{RT_0}$$

$$\text{Hydrostatic eq.: } dp = -\rho g dz$$

$$\Rightarrow dp = -\frac{p}{RT_0} g dz$$

$$\Rightarrow \frac{dp}{p} = -\frac{g}{RT_0} dz$$

$$\Rightarrow \int_{\rho(z=0)=\rho_0}^{\rho(z)} \frac{dp}{p} = \int_{z=0}^z \left(-\frac{g}{RT_0}\right) dz$$

$$\Rightarrow \ln\left(\frac{\rho(z)}{\rho_0}\right) = -\frac{g}{RT_0} z = -\frac{z}{H}, \quad H = \frac{RT_0}{g}$$

$$\Rightarrow \boxed{\rho(z) = \rho_0 e^{-\frac{z}{H}}}, \quad H = \frac{RT_0}{g} = \text{"scale height"}$$

(b) Polytropic atmosphere:  $T(z) = T_0 - \gamma z$ ,  
 $T_0, \gamma$  const.

$$\text{Gas eq. + hydr. eq. } \Rightarrow dp = -\frac{\rho g}{R T(z)} dz$$

$$\Rightarrow \frac{dp}{p} = -\frac{g}{R(T_0 - \gamma z)} dz$$

Extra-credit problem (cont'd)

Substitution:  $y = T_0 - \gamma z$ ,  $\frac{dy}{dz} = -\gamma$ ,  $dz = -\frac{1}{\gamma} dy$

$$\Rightarrow \frac{dp}{p} = -\frac{\gamma}{R\gamma} \left(-\frac{1}{\gamma} dy\right) = \frac{\gamma}{R\gamma} \frac{dy}{y}$$

$$\Rightarrow \int_{p_0}^{p(z)} \frac{dp}{p} = \int_{y_0}^{y(z)} \frac{\gamma}{R\gamma} \frac{dy}{y}$$

$$\Rightarrow \ln \frac{p(z)}{p_0} = \frac{\gamma}{R\gamma} \ln \frac{y(z)}{y_0} = \frac{\gamma}{R\gamma} \ln \frac{T_0 - \gamma z}{T_0}$$

$$\Rightarrow \frac{p(z)}{p_0} = e^{\frac{\gamma}{R\gamma} \ln \frac{T_0 - \gamma z}{T_0}} = \left( e^{\ln \frac{T_0 - \gamma z}{T_0}} \right)^{\frac{\gamma}{R\gamma}}$$

$$\Rightarrow \boxed{p(z) = p_0 \left( \frac{T_0 - \gamma z}{T_0} \right)^{\frac{\gamma}{R\gamma}}}$$

(c) Recall:  $e = \lim_{x \rightarrow \infty} \left[ \left( 1 + \frac{1}{x} \right)^x \right]$

Here:  $\frac{T_0 - \gamma z}{T_0} = 1 - \frac{\gamma z}{T_0}$ , so let  $\frac{1}{x} = -\frac{\gamma z}{T_0}$

Extra-credit problem (cont'd)

$$\left( \begin{array}{l} \hookrightarrow \\ \end{array} \right. \quad x = -\frac{T_0}{yz} \quad , \quad \frac{1}{y} = -\frac{z}{T_0} x$$

$$\lim_{y \rightarrow 0} \left( 1 - \frac{yz}{T_0} \right)^{\frac{g}{R} y}$$

$$= \left( 1 + \frac{1}{x} \right)^{\frac{g}{R} \left( -\frac{z}{T_0} x \right)}$$

$$= \left[ \left( 1 + \frac{1}{x} \right)^x \right]^{-\frac{g}{RT_0} z}$$

$$\rightarrow \lim_{y \rightarrow 0} p(z) = \lim_{x \rightarrow \infty} \left( p_0 \left[ \left( 1 + \frac{1}{x} \right)^x \right] \right)^{-\frac{g}{RT_0} z}$$

$$= p_0 \cdot \left[ \lim_{x \rightarrow \infty} \left( 1 + \frac{1}{x} \right)^x \right]^{-\frac{g}{RT_0} z} = p_0 e^{-\frac{g}{RT_0} z}$$

q.e.d.