

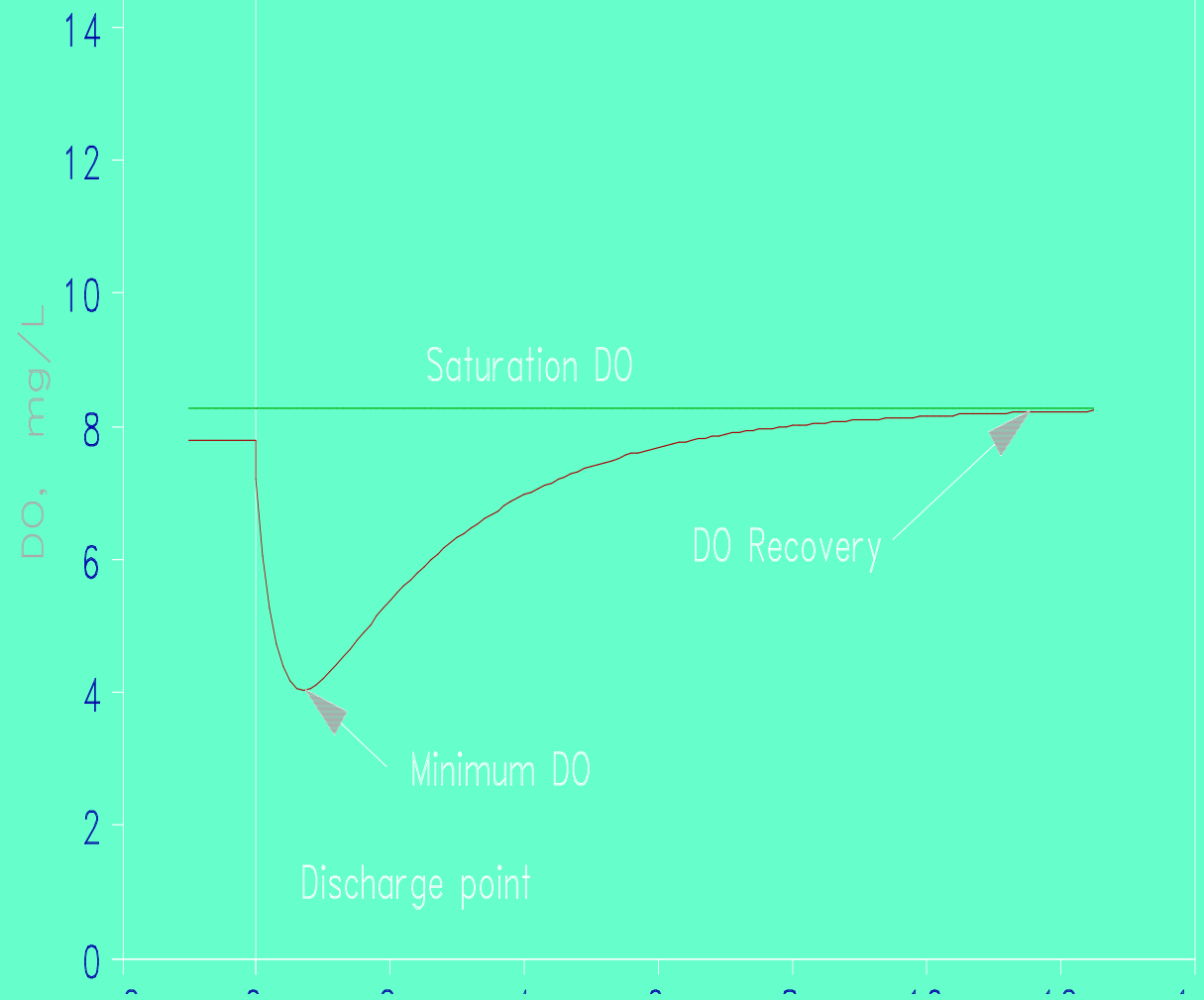
# CEE 577: Surface Water Quality Modeling

Lecture #16

Streeter-Phelps: Reaeration & Dams

(Chapra, L22 & L23)

# Typical DO Sag Curve



# Streeter Phelps Equation

$$V \frac{dD}{dt} = k_d VL - k_a VD$$

$$L = L_o e^{-k_d t}$$

This equation can be solved by separation of variables and integration, or by use of an integrating factor. The boundary condition is  $t = 0 @ D = D_o$ . This yields the DO sag

$$D = D_o e^{-k_a t} + \frac{k_d L_o}{k_a - k_r} \left( e^{-k_r t} - e^{-k_a t} \right)$$

where

$D$  = stream deficit at time  $t$ , [mg/L]

$D_o$  = initial oxygen deficit (@  $t = 0$ ), [mg/L]

And recognizing that:  $t=x/U$

$$D = D_o e^{-k_a x/U} + \frac{k_d L_o}{k_a - k_r} \left( e^{-k_r x/U} - e^{-k_a x/U} \right)$$

Anaerobic Conditions?

# Critical Time

- The most stress is placed on the aquatic life in a stream when the DO is at a minimum, or the deficit,  $D$ , is a maximum. This occurs when  $dD/dt = 0$ . We can obtain the time at which the deficit is a maximum by taking the derivative of the DO sag equation with respect to  $t$  and setting it equal to zero, then solving for  $t$ . This yields,

$$t_{crit} = \frac{1}{k_a - k_r} \ln \left[ \frac{k_a}{k_r} \left( 1 - \frac{D_o(k_a - k_r)}{k_d L_o} \right) \right]$$

$t_{crit}$  = time at which maximum deficit (minimum DO) occurs, [days]

# Special Case for $D$ and $t_c$

- When  $k_a$  and  $k_r$  are equal:

$$t_c = \frac{1}{k_r} \left( 1 - \frac{D_0}{L_0} \right)$$

$$D = (L_0 k_r t + D_0) e^{-k_r t}$$

From Davis & Masten, page 290-291

# Critical concentration

- Once the critical time is known, you can calculate the  $c_{\min}$

$$c_{\min} = c_s - L_o \frac{k_d}{k_a} e^{-k_r t_{\text{crit}}}$$

- this is an abbreviated form of the full equation, which is only valid for  $t_{\text{crit}}$

**This differs from Chapra's equation 21.14 on page 397,  
Why??????**

# Estimating Reaeration Rates ( $d^{-1}$ )

(U in ft/s; H in ft)

- O'Connor-Dobbins formula
  - based on theory
  - verified with some deep waters

$$k_a = 12.9 \frac{U^{0.5}}{H^{1.5}}$$

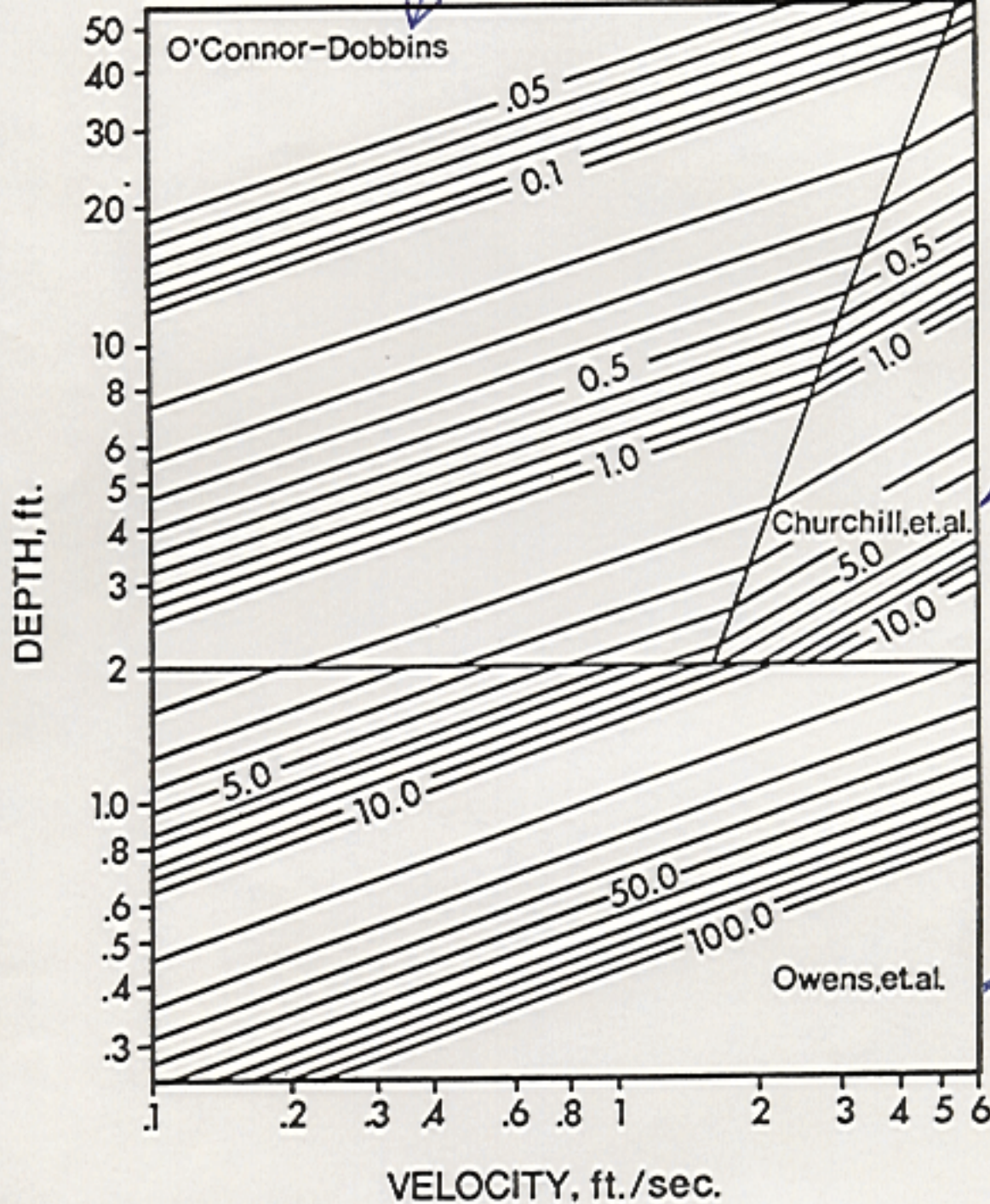
- Churchill formula
  - Tennessee Valley
  - deep, fast moving streams

$$k_a = 11.6 \frac{U}{H^{1.67}}$$

- Owens formula
  - British
  - shallow streams

$$k_a = 21.6 \frac{U^{0.67}}{H^{1.85}}$$





## The method of Covar (1976)

**Values for  $k_a$  are in units of  $d^{-1}$**

# Using the Covar approach

- Determine proper Domain

H (ft)	U (ft/s)	Formula
<2	Any	Owens
>2	$<1.2H^{0.34}$	O'Connor-Dobbins
>2	$>1.2H^{0.34}$	Churchill

# More on $k_a$ estimation

- Tsivoglou & Wallace (1972) method
  - $k_a = 0.88US$ , for  $Q = 10-300$  cfs
  - $k_a = 1.8US$ , for  $Q = 1-10$  cfs
- Temperature correction
  - $\theta = 1.024$

$$k_T = k_{20^\circ C} \theta^{T-20^\circ C}$$

# Dam Reaeration

- Butts and Evans (1983):

$$r = 1 + 0.38abH(1 - 0.11H)(1 + 0.046T)$$

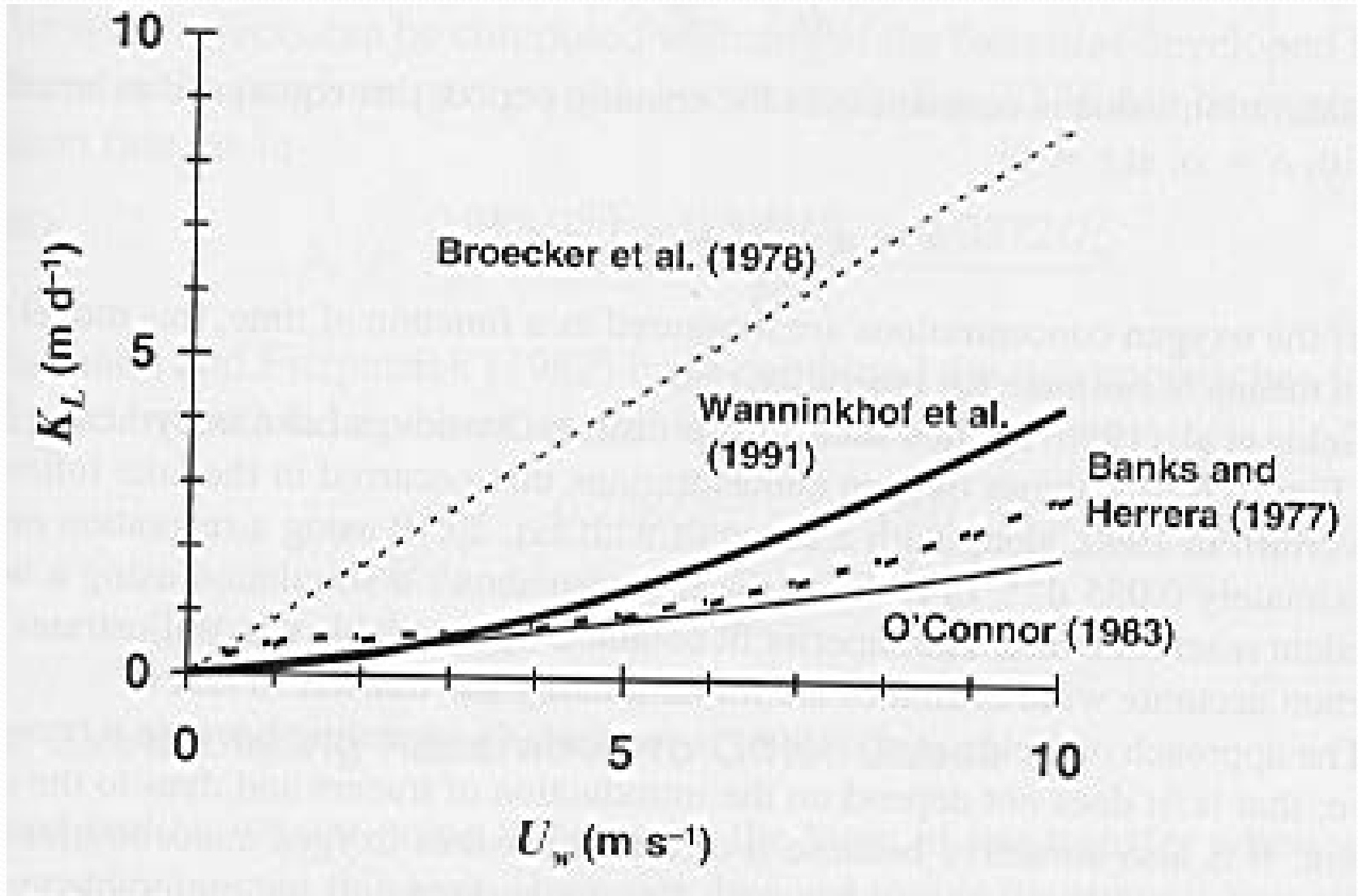
Temperature (°C)

Ratio of deficit  
above and  
below the dam

Empirical  
coefficients  
which relate to  
water quality and  
dam type (Table  
20.2)

Difference in water  
elevation

# Wind dependent reaeration formulas



- To next lecture